

An Ocean Heat Budget for the Florida Reef Tract

Sea temperature variability, air-sea
flux, and the *thermal siphon*



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Significance of reef sea temperature

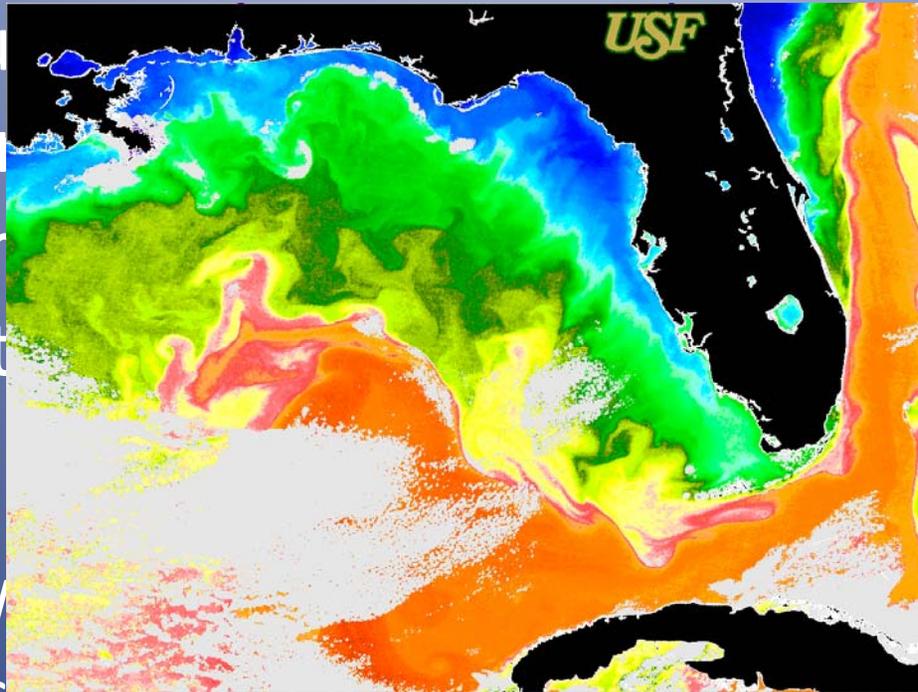
- Thermal environment
 - Sea temperature is a control on growth rate of reef-building corals (e.g., Cantin et al. 2010)
 - Extremes can cause **bleaching** and mortality
 - Can play a role in spawning & settlement success
- As a proxy for studying nutrient fluxes
 - **Upwelling** (T/N,P relationships below thermocline, e.g., Leichter et al. 2003, Hitchcock et al. 2005)
 - Cross-shore flows (land sources)

Reef Sea Temperature and **Climate**

- Our current understanding of the global air-sea climate system comes from coupled numerical models at **regional** scales, using satellite data at **kilometer** scales and greater.
- To downscale long-term forecasts to coral reef ecosystem impacts means accounting for physical processes at all scales – from global systems to *individual coral communities*.

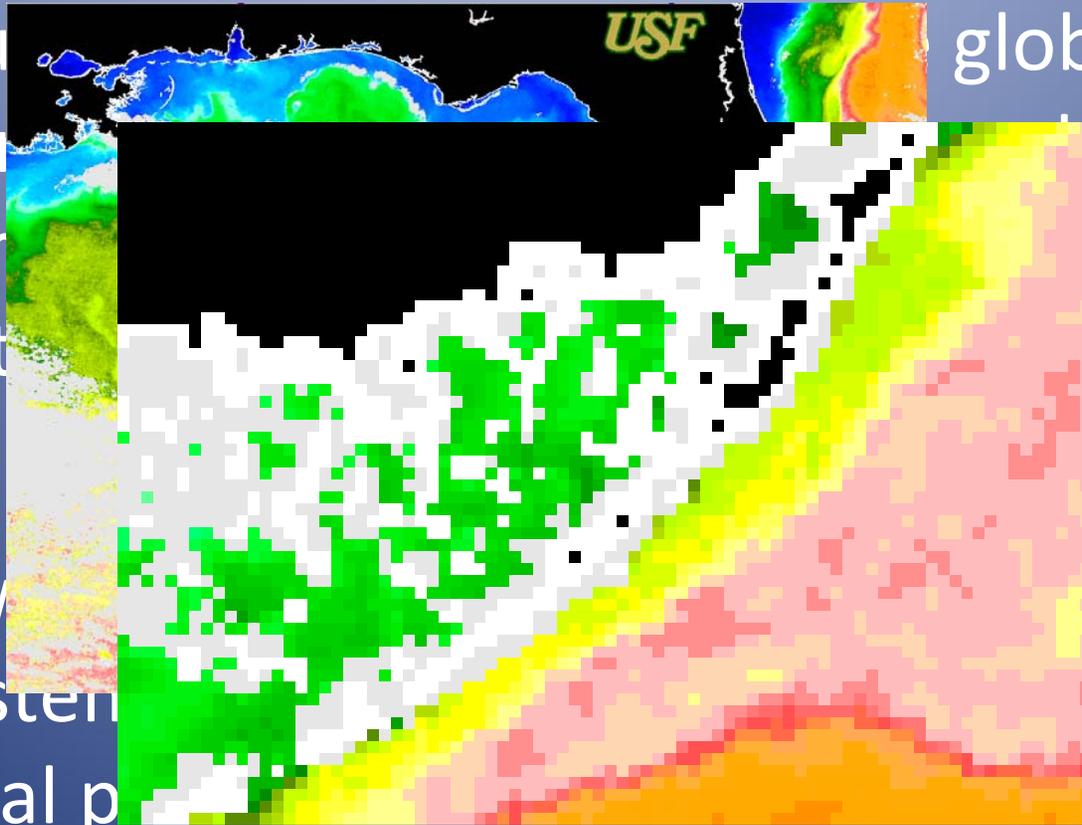
Reef Sea Temperature and **Climate**

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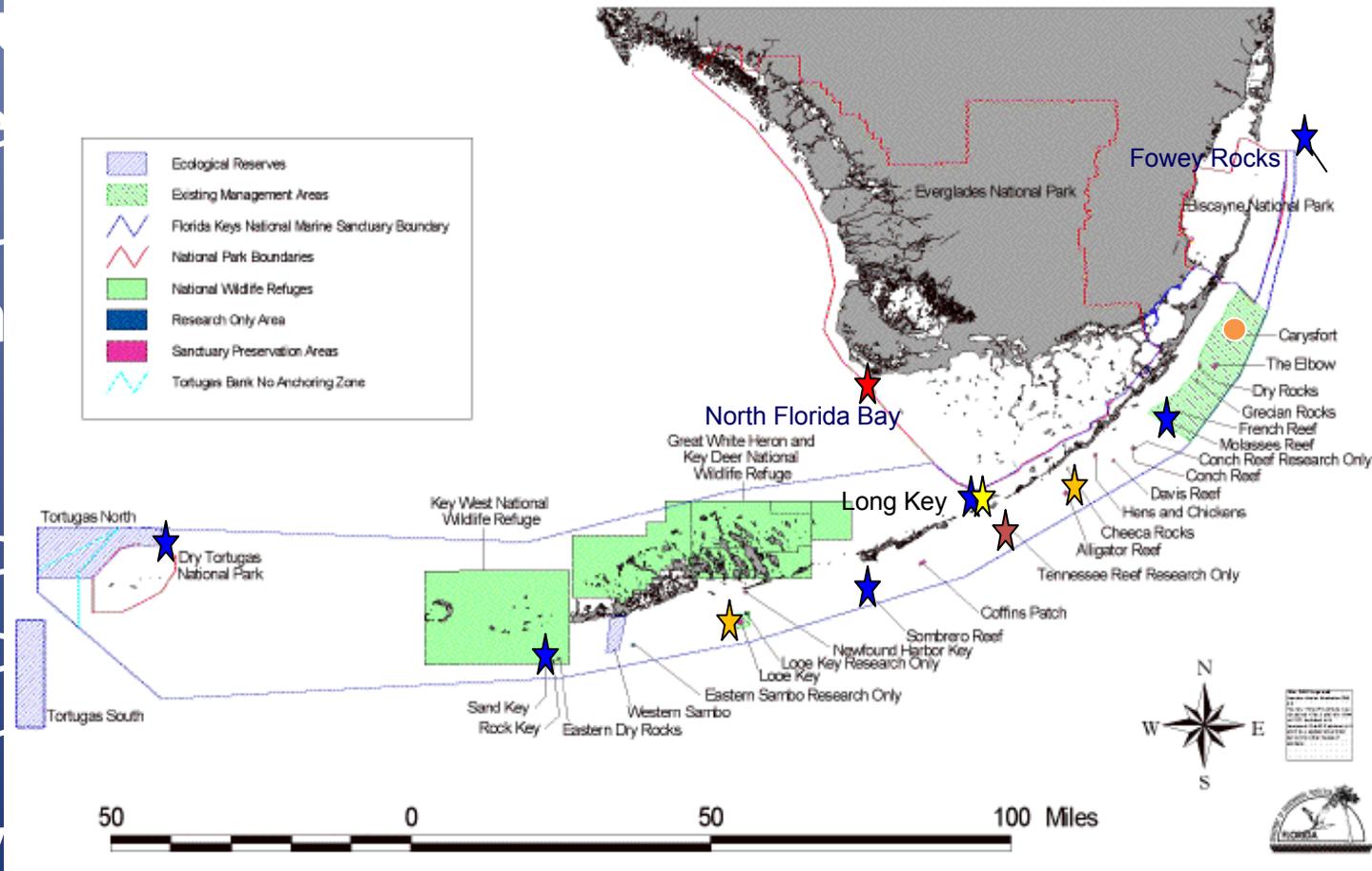
Reef Sea Temperature and Climate

- Our current understanding of reef sea climate is based on a limited number of numerical models and satellite data. global air-sea climate is changing and becoming more extreme.
- To develop a better understanding of reef sea climate, we need to move from global systems to individual coral communities.



Reef Sea Temperature and Climate

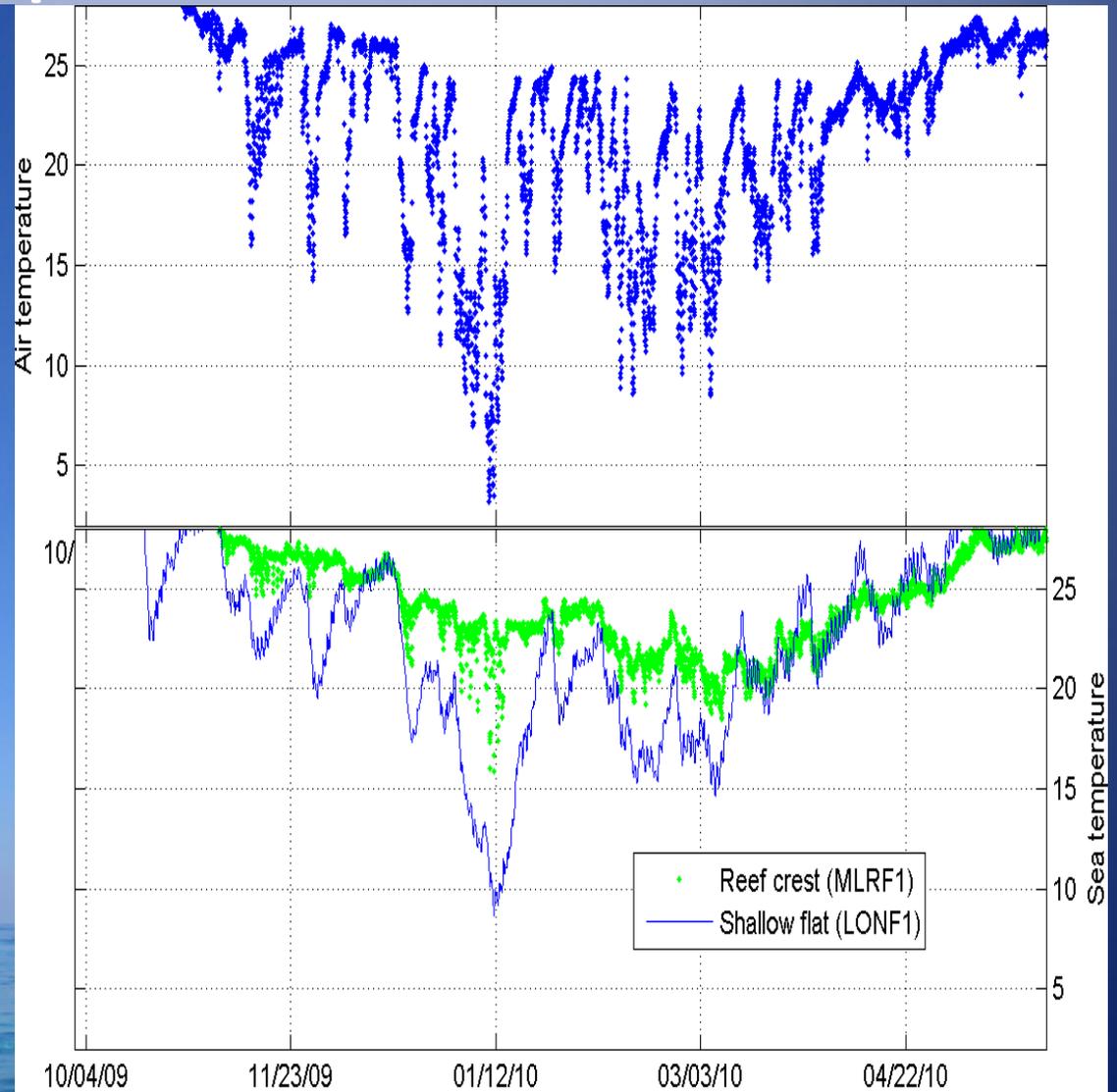
Florida Keys National Marine Sanctuary



- Oceanic
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- sy

• Reef

Reef Sea Temperature and Climate



Reef Sea Temperature **Variability**

- Rapid cooling events in summer can often be identified with storm-induced surface waves and **storm surge** (Manzello et al. 2007)
- Higher than normal variability can also be coincident with passage of **eddies** offshore (Gramer et al. 2009), and with shoaling of **internal waves** (e.g., Davis et al. 2008).
- To characterize these events in the record, we must account for all sources of ocean heating.

Heat budget

$$T_t = -\vec{u}_{km} \cdot \nabla_h T_{km} - \vec{u}_{ge} \cdot \nabla_h T_{sfc} - \vec{u}_{hc} \cdot \nabla_h T_{hc} (Q_0, h, \beta) + \frac{Q_0}{\rho C_p h}$$

- Air-sea fluxes:
 - Turbulent – TOGA-COARE 3.0a (Fairall et al. 2003)
 - Radiative – NCEP NARR (32km reanalysis, Mesinger et al. 2006)
- Kilometer-scale advection: sources of estimates
 - Gulf of Mexico HYCOM (4km res., 2003-2010)
 - Florida Keys HYCOM (~900m res., 2008)
- Sub-kilometer scale advection
(Only parameterized in kilometer-scale ocean models)
 - Wind- and swell-driven surface transport? (Ardhuin et al. 2009)

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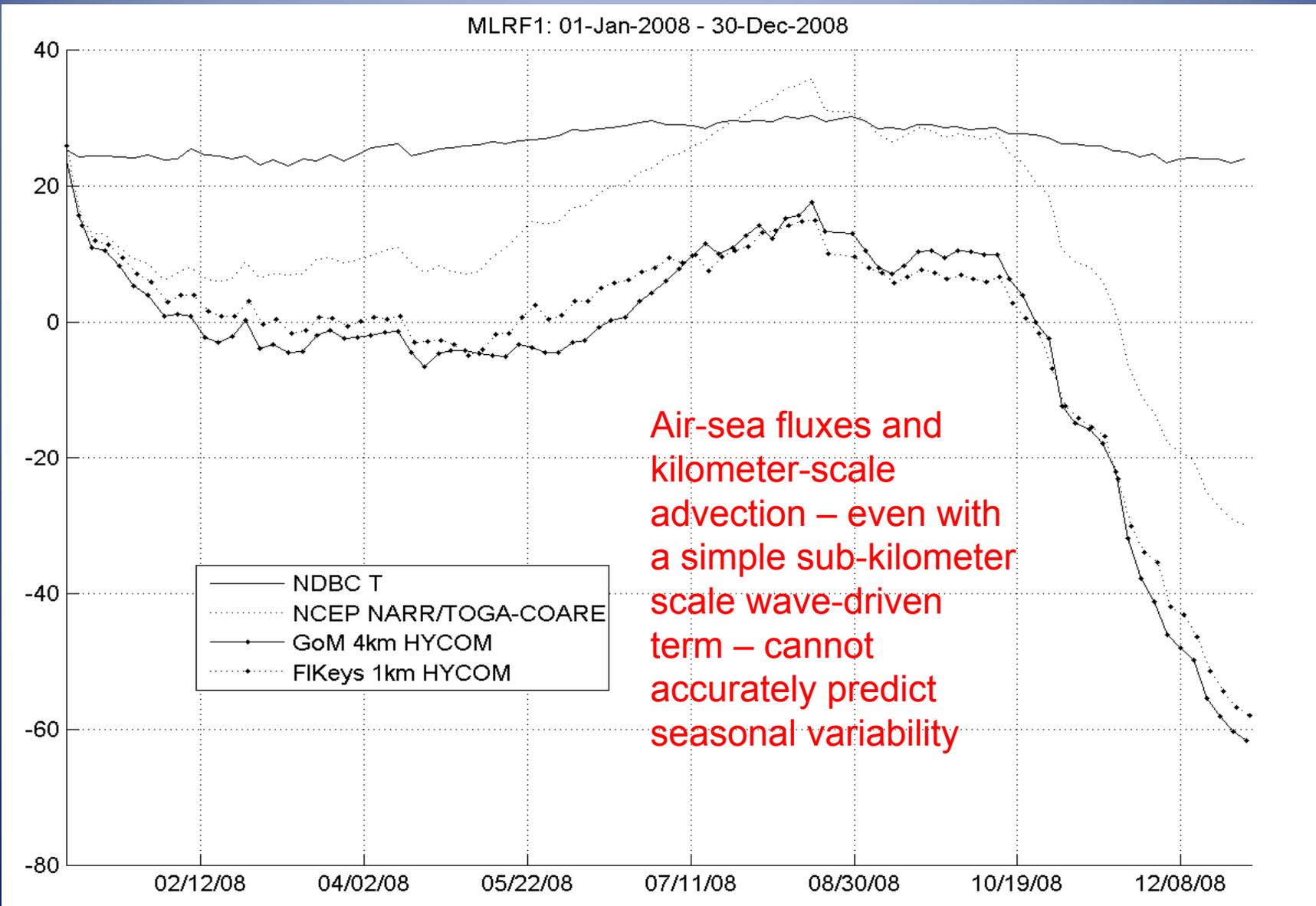
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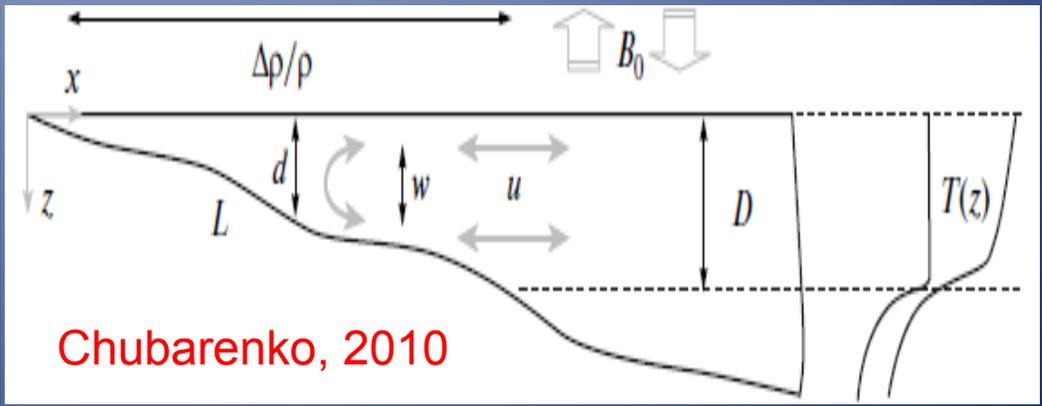
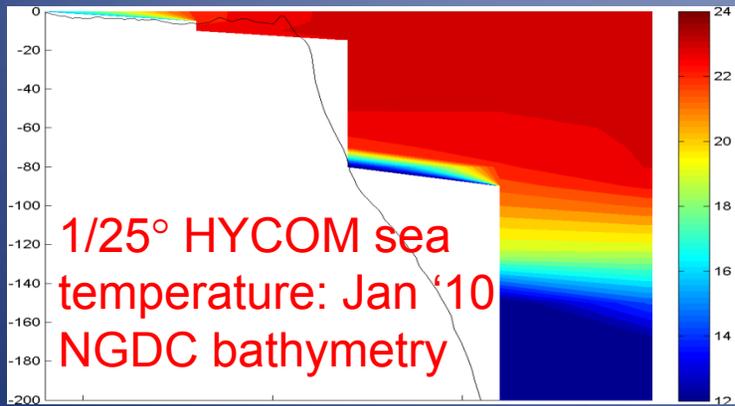
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Results – air-sea and km-scale only



Horizontal Convection – *thermal siphon*

- Air-sea flux and km-scale heat advection alone do not model variability well: is there also a smaller-scale oceanographic process at work?

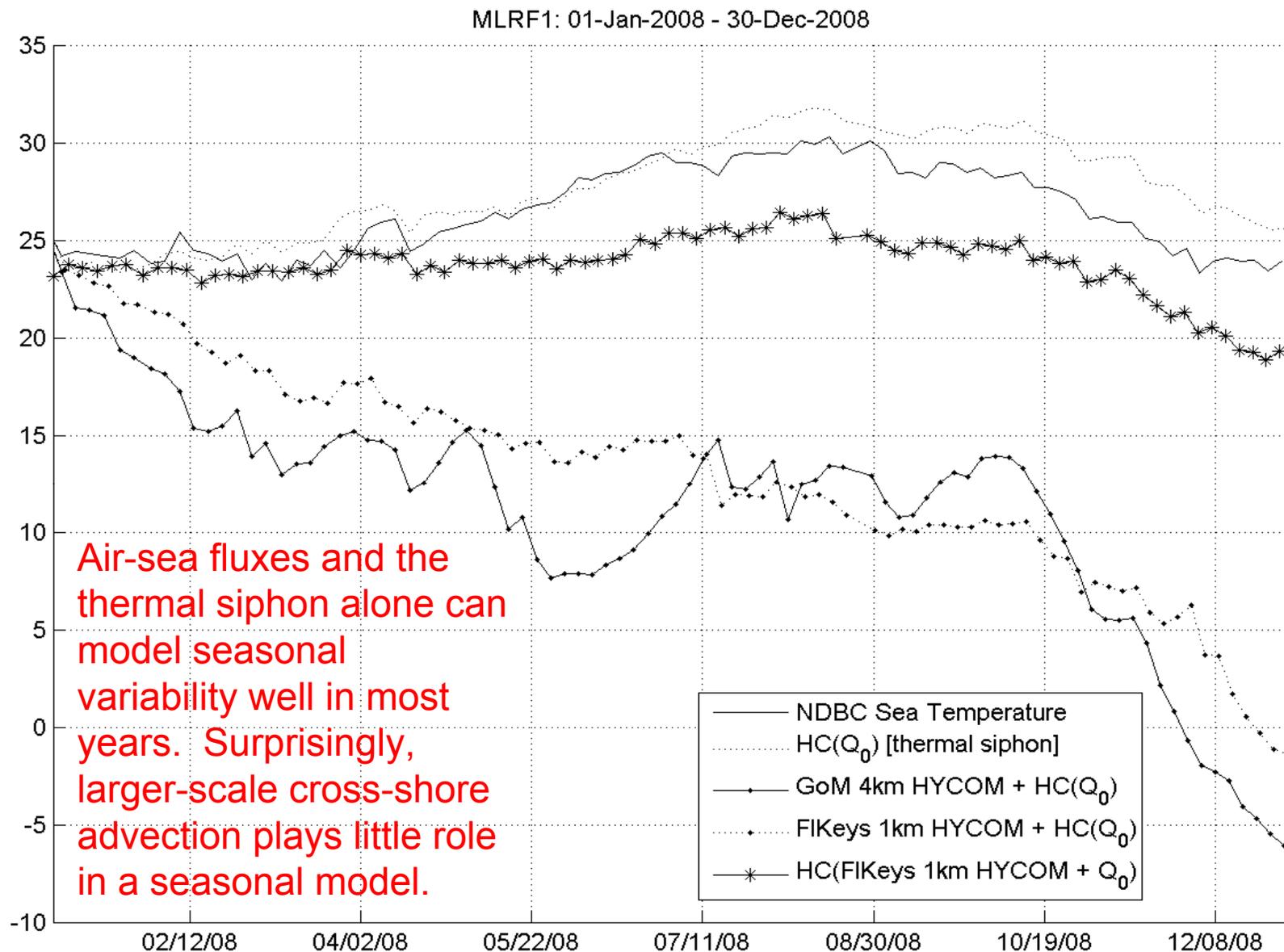


Monismith et al., 2006

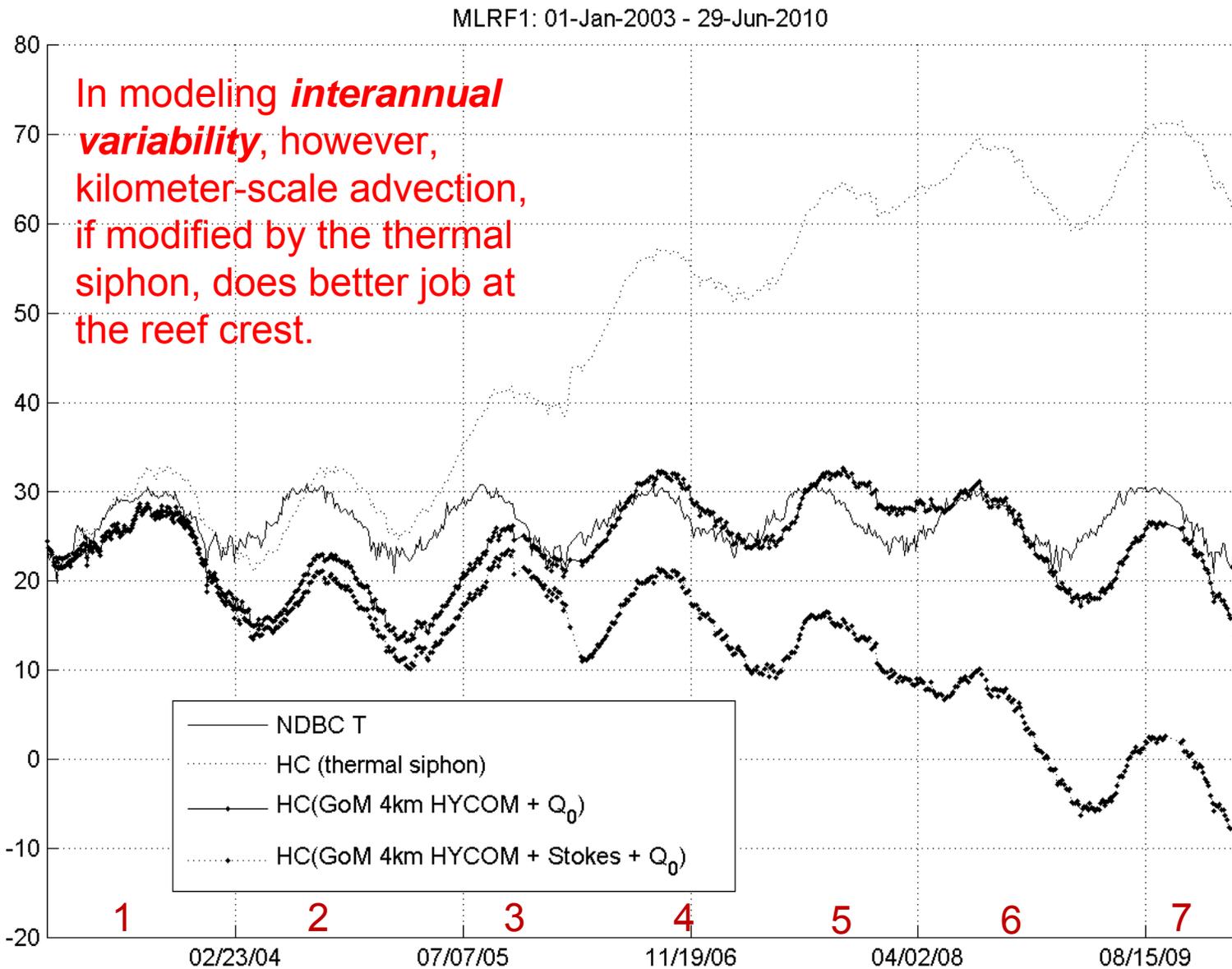
TABLE 1. Possible scalings for ΔV .

Balance	Unsteady inertia	Stress divergence	Advective inertia
Steady temperature	$V_{US} \sim (u_f^3 T / D)^{1/2}$	$V_{VS} \sim (u_f^{3/2} / q^{1/2})$	$V_{NS} \sim \beta^{-1/3} u_f$
Unsteady temperature	$V_{UU} \sim (\beta u_f^3 T^2 / D_0^2)$	$V_{VU} \sim (\beta u_f^3 T / q D)$	$V_{NU} \sim (u_f^3 T / D)^{1/2}$

Results (various models, with and without siphon)

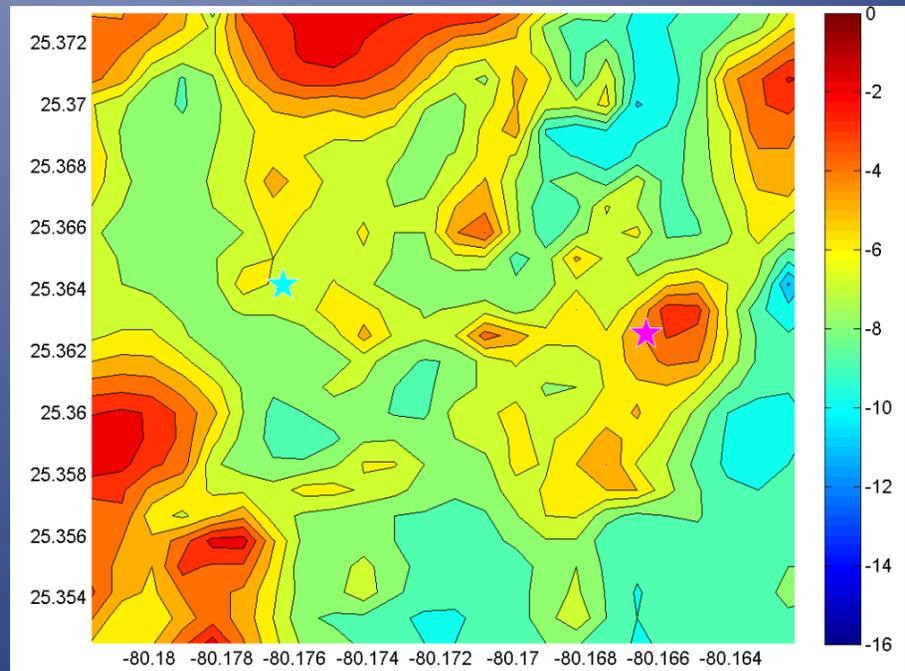
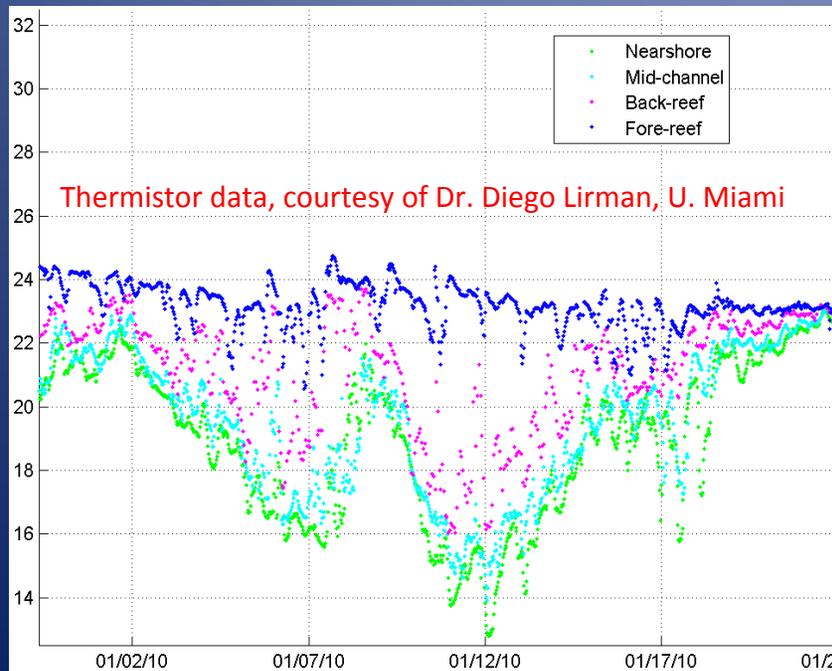


Results – reef crest sites



Results – low-relief areas

- Two Upper Keys thermistor sites at similar depths of 4-5m, one on a mid-Channel flat, the other a back-reef promontory (magenta and cyan stars in bathymetry map below): in the cold snap of 2010, the same air-sea forcing brought very different responses between the two sites – the thermal siphon only being able to moderate sea temperature variability effectively at the higher-relief site (magenta).



Future Work

- Apply heat budget to analyze higher-frequency (*weekly to diurnal period*) variability.
- Model extreme events (e.g., *upwelling*) not well explained by the current heat budget
- Adapt the heat budget as a *management tool*, recognizing patterns in physical data likely to impact effective management of reef ecology.

Acknowledgements, References

In situ data from NOAA NDBC and FIO, and Dr. Diego Lirman, UM RSMAS; insightful comments by Drs. Derek Manzello and Kevin Helmle, NOAA AOML

Reanalysis and model data from: NOAA National Centers for Environmental Prediction (NCEP), Naval Research Lab / HYCOM Consortium, Dr. V. Kourafalou-RSMAS, NOAA Wave Watch III®

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THANK YOU!

Photo © J. E. N. Veron, Corals of the World, 2000
Large *Porites astreoides* ("mustard hill coral")

<http://ecoforecast.coral.noaa.gov>

Additional Slides

Background - SEAKEYS

– C-MAN / SEAKEYS monitoring stations

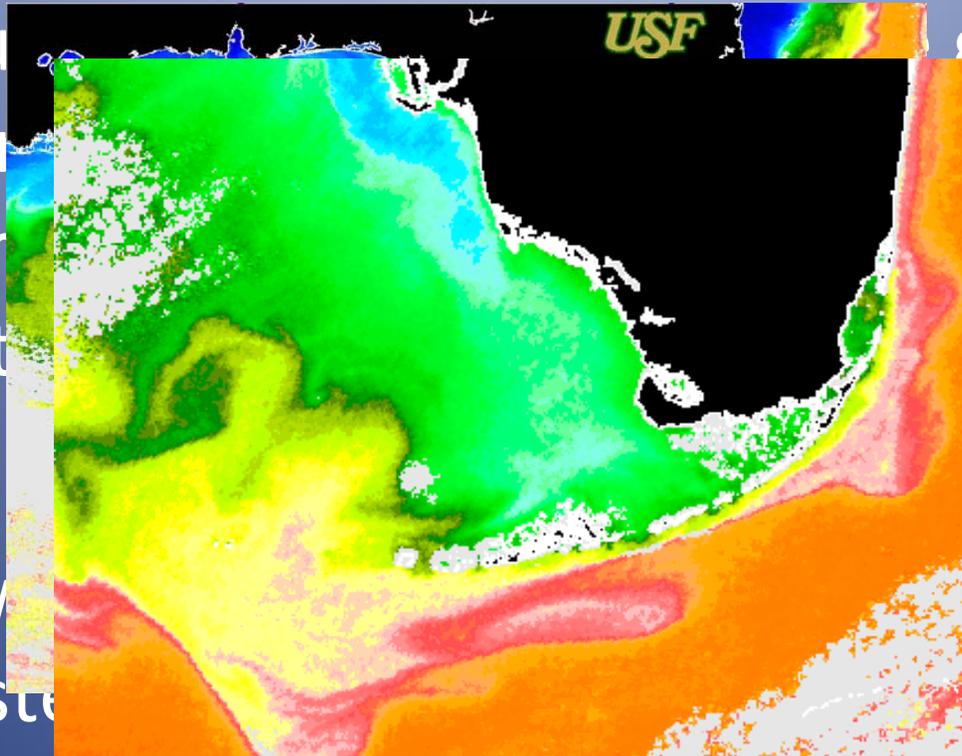
- At sites on the reef crest (8km offshore) and near shore
- Continuous long-term hourly records (>20 years)
- Sea temperature, meteorology, and in some cases salinity, tide depth, surface and sub-surface light, and other variables
- Quality control procedures applied to most variables

– Hourly sea temperature record

- Energetic variability at multiple periodicities: 6- and 8-hour, M_2 tidal, diurnal, inertial (28h), seasonal, *interannual trends*
- Strong mean squared coherence with air temperature at diurnal, seasonal, 4-, 12-, and 80-day periodicities
- Occasional periods of very low diurnal coherence

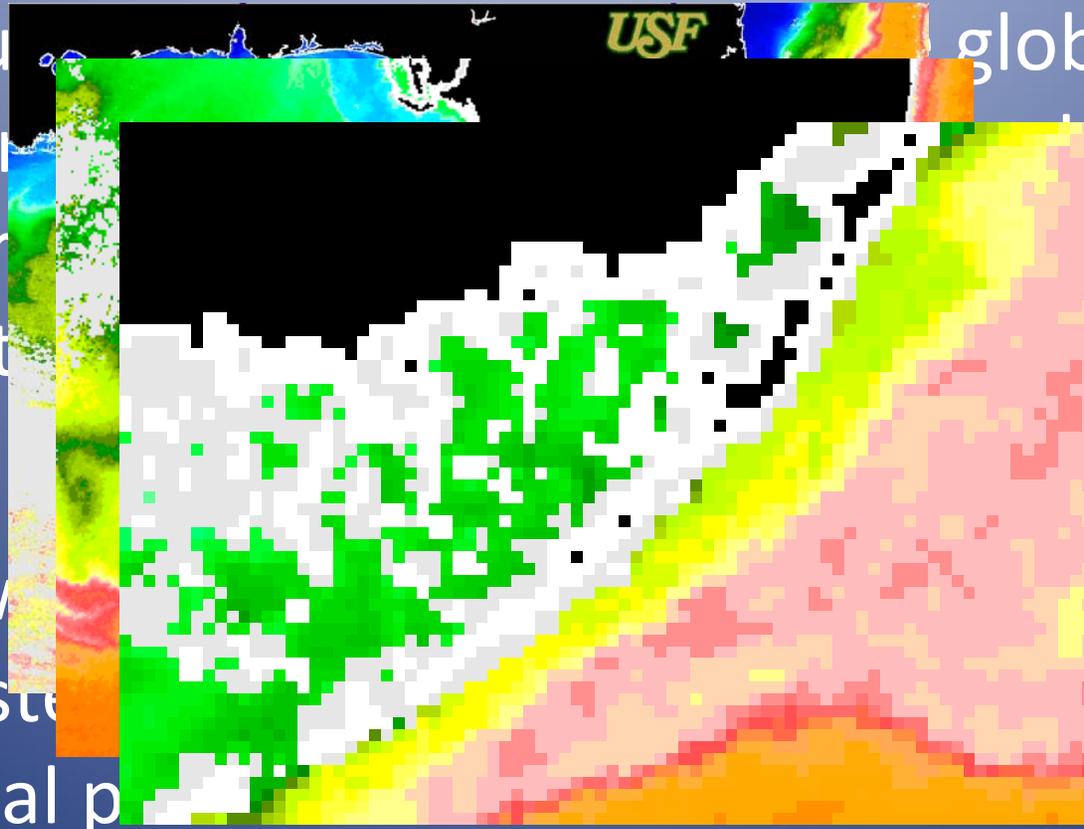
Reef Sea Temperature and Climate

- Our current understanding of reef sea climate is based on a combination of global air-sea climate coupled models, using satellite data and greater. To downscale these models to coral reef ecosystems, we are developing physical processes at all scales – from global systems to individual coral communities.

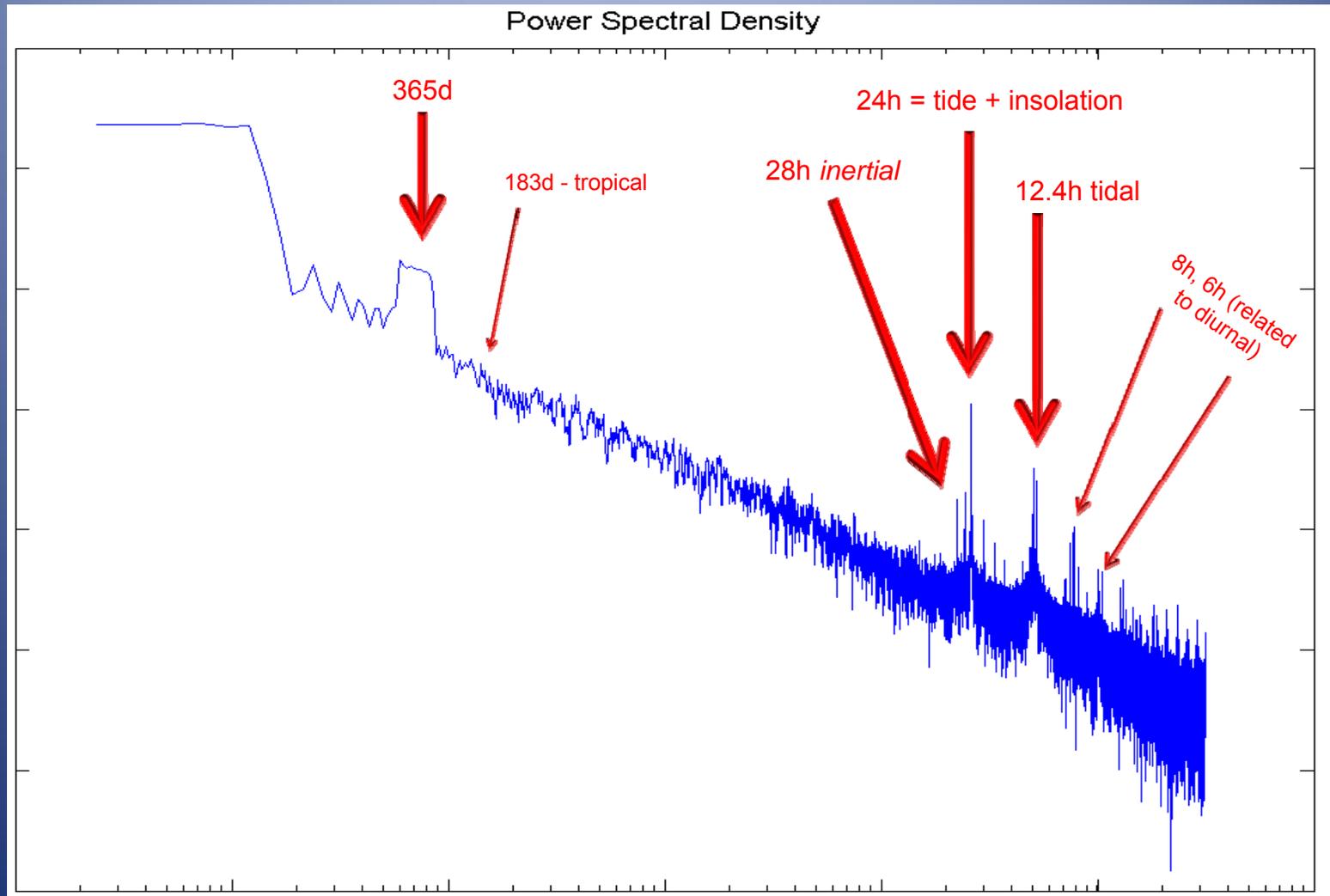


Reef Sea Temperature and Climate

- Our current understanding of global air-sea climate is based on a limited number of observations, and satellite data is becoming increasingly available. To do this, we need to move from global systems to individual coral communities.

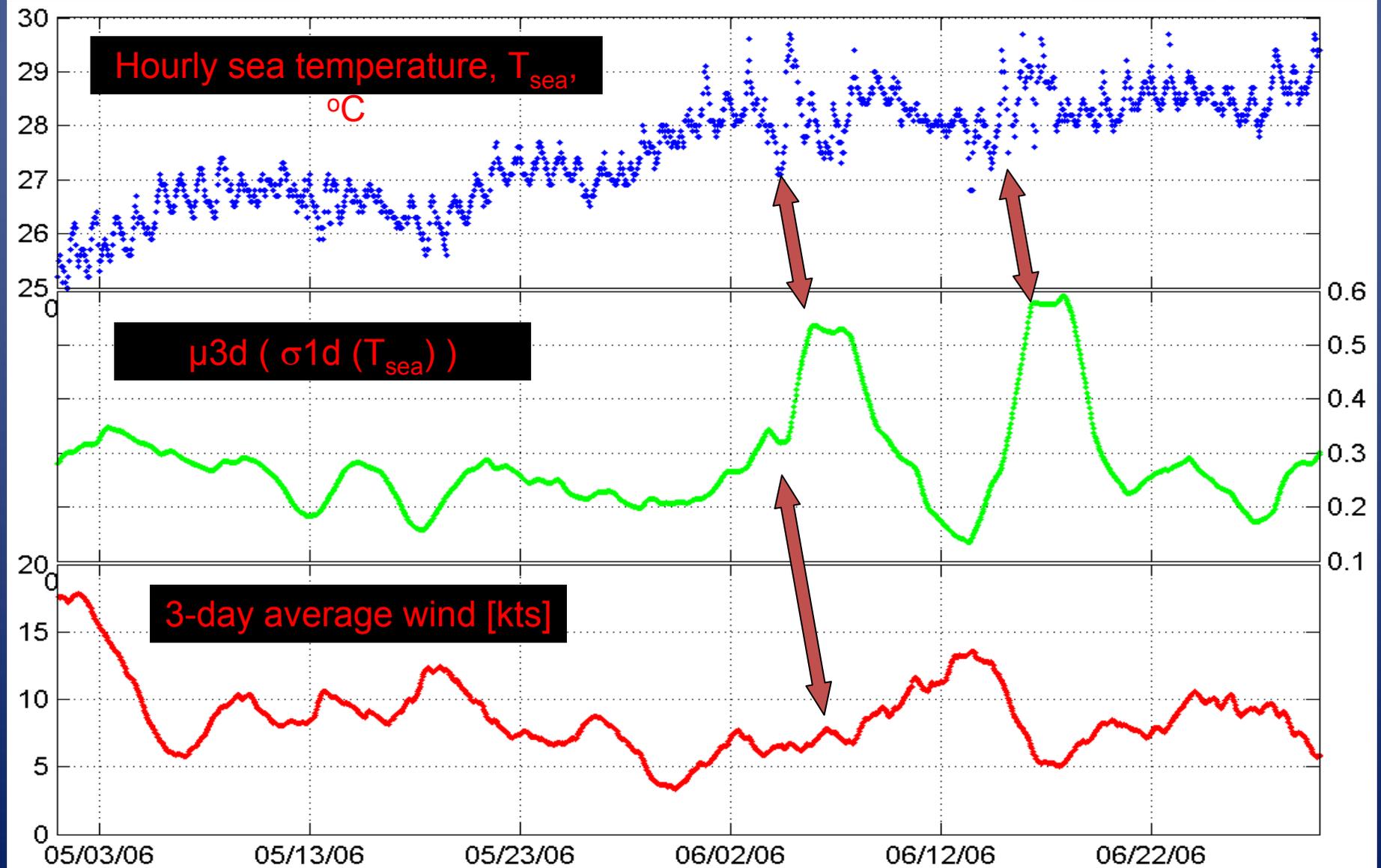


Sea temperature spectrum

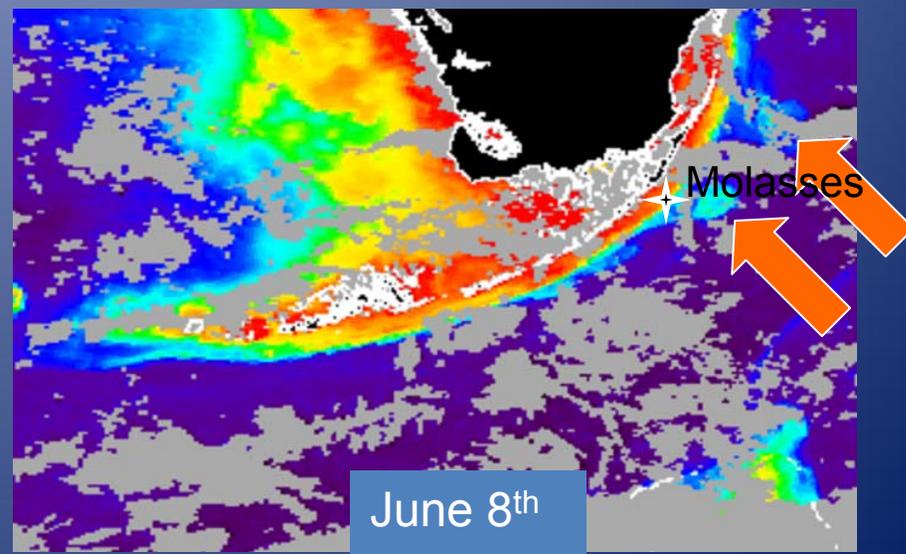
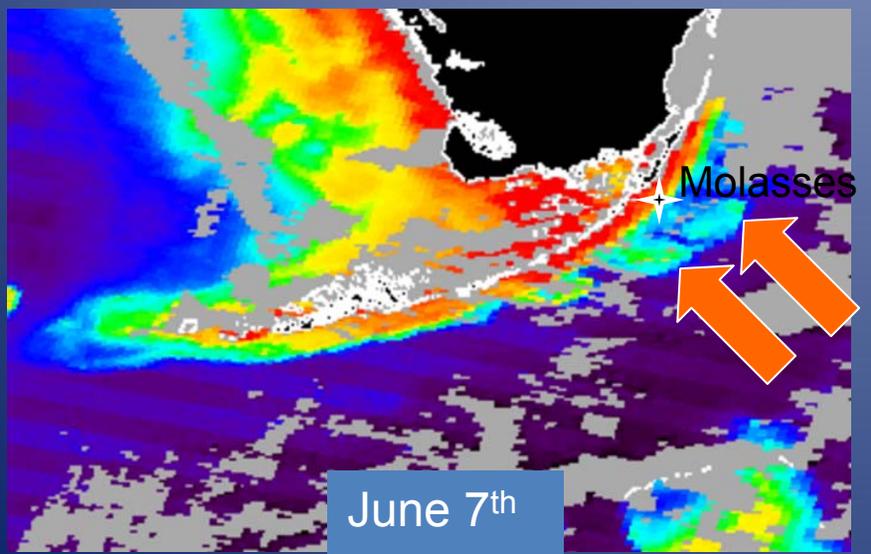
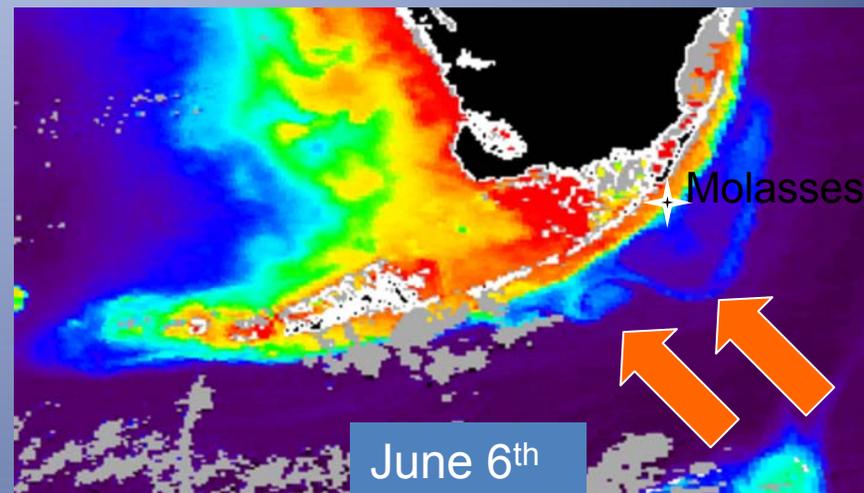
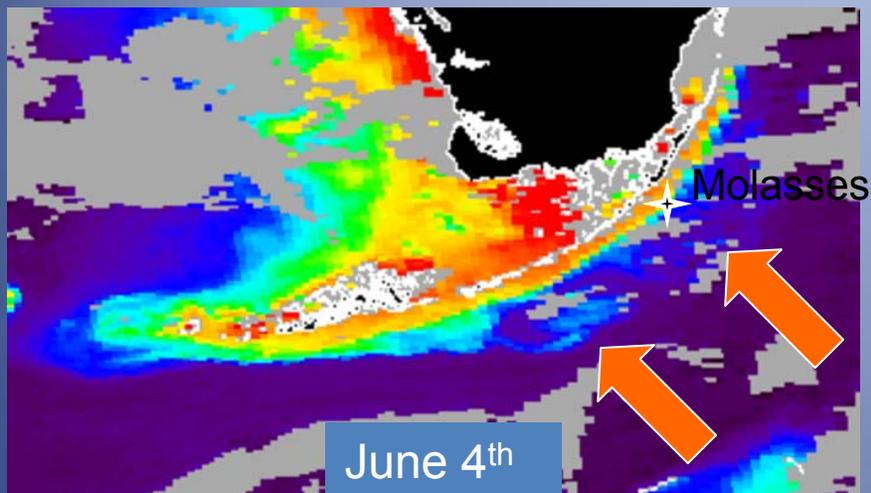


Oceanographic and air-sea processes at a range of time/space scale

Initial Motivation – June 2006 example



June 2006 – *cyclonic* chlorophyll features



Heat budget

- Sea temperature variability
 - How much is forced by direct air-sea flux?
 - Horizontal mixing and advection? Vertical mixing and upwelling (winds, eddies, internal waves)?

$$T_t = -\vec{u}_{km} \cdot \nabla_h T_{km} - \vec{u}_{qe} \cdot \nabla_h T_{sfc} - \vec{u}_{hc} \cdot \nabla_h T_{hc} (Q_0, h, \beta) + \frac{Q_0}{\rho C_p h}$$

$$Q_0 = \gamma Q_{SW} + Q_{LW} + Q_{SH} + Q_{LH} + Q_{RH}$$

$$\gamma = 1 - A_b \cdot \tau_{PAR} \cdot (1 - \tau_{PAR}), \quad \tau_{PAR} = \exp(-K_d^{PAR} \cdot h \cdot \sec(\theta))$$

$$|\vec{u}_{qe}| \cong 5.0 \times 10^{-4} \left[1.25 - 0.25 \left(\frac{0.5}{f_c} \right)^{1.3} \right] |U_{10} \times \min\{U_{10}, 14.5\}| + 0.025(H_s - 0.4)$$

$$Q_{hc}^V = C \cdot \beta \cdot u_f^3 \cdot T_f^2 / h^2$$

$$u_f = (\alpha \cdot g \cdot h \cdot H_0)^{1/3},$$

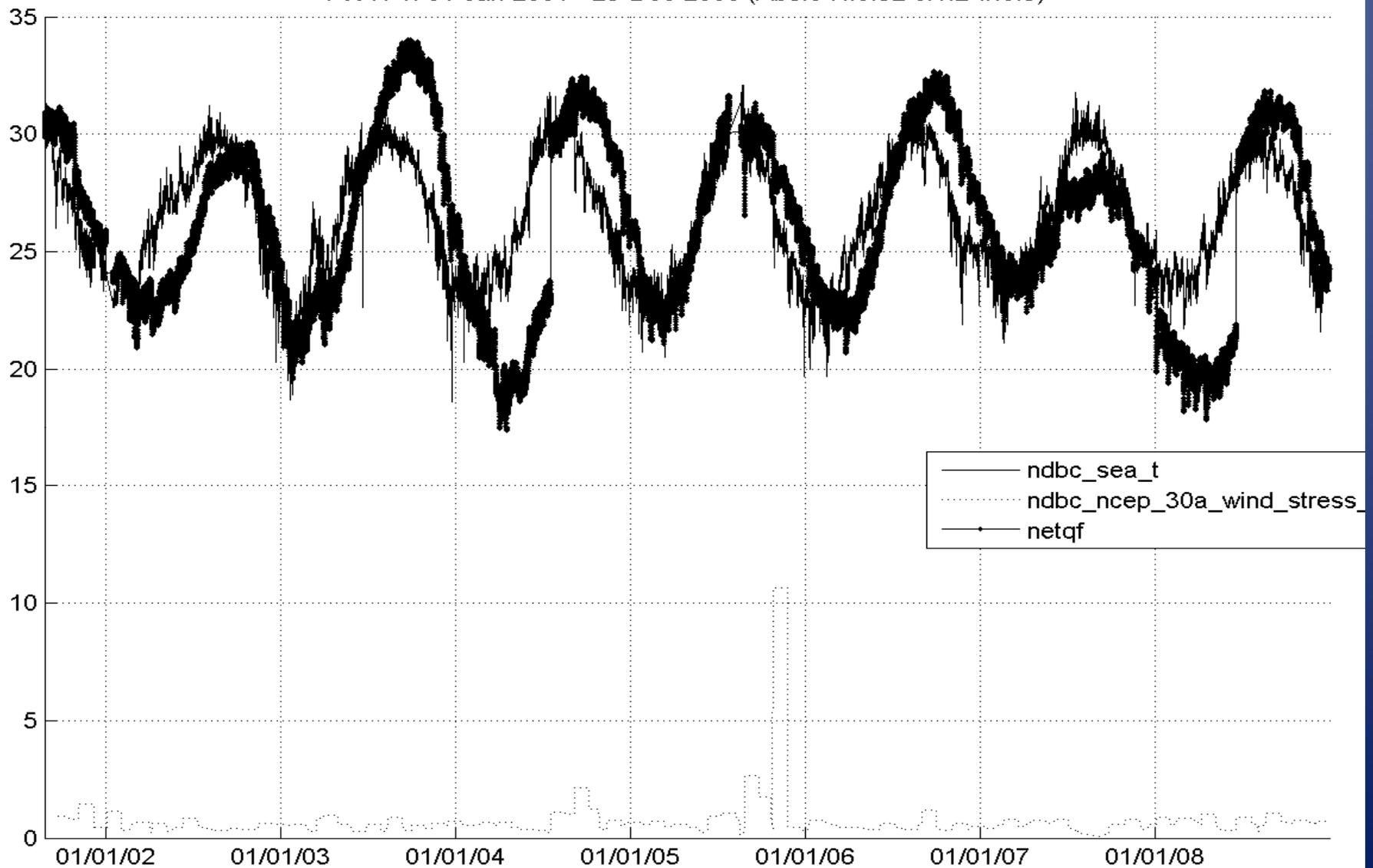
$$H_0 = Q_{24hSMA} / \rho C_p \cdot$$

$$\bar{u}_{hc} = \frac{Q_{hc}^V}{h};$$

$$|\nabla_h T_{hc}| = \frac{H_0}{\rho C_p} \left(\frac{1}{h} - \frac{1}{(h + \beta l)} \right)$$

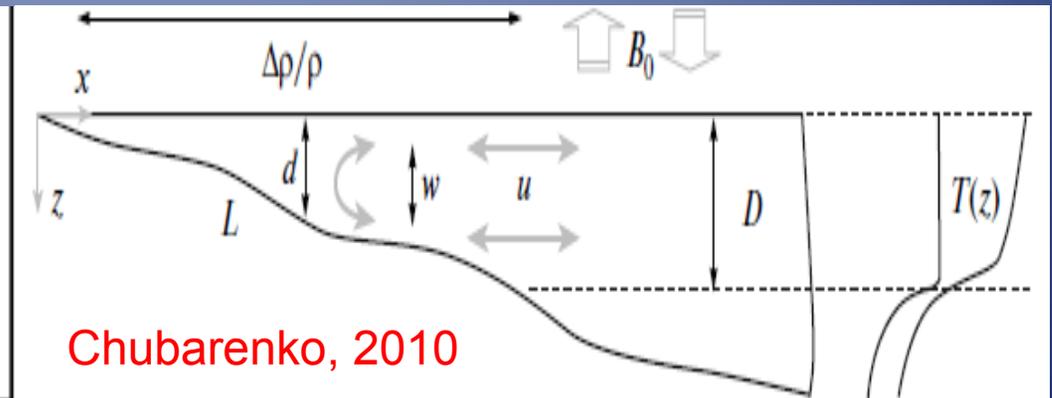
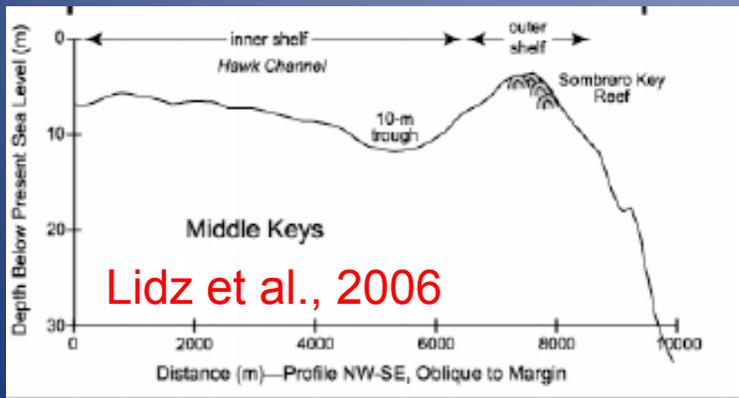
Heat Budget - Results

FWYF1: 01-Jan-2001 - 29-Dec-2008 (Abs:0 R:0.92 c:1.2 w:0.3)



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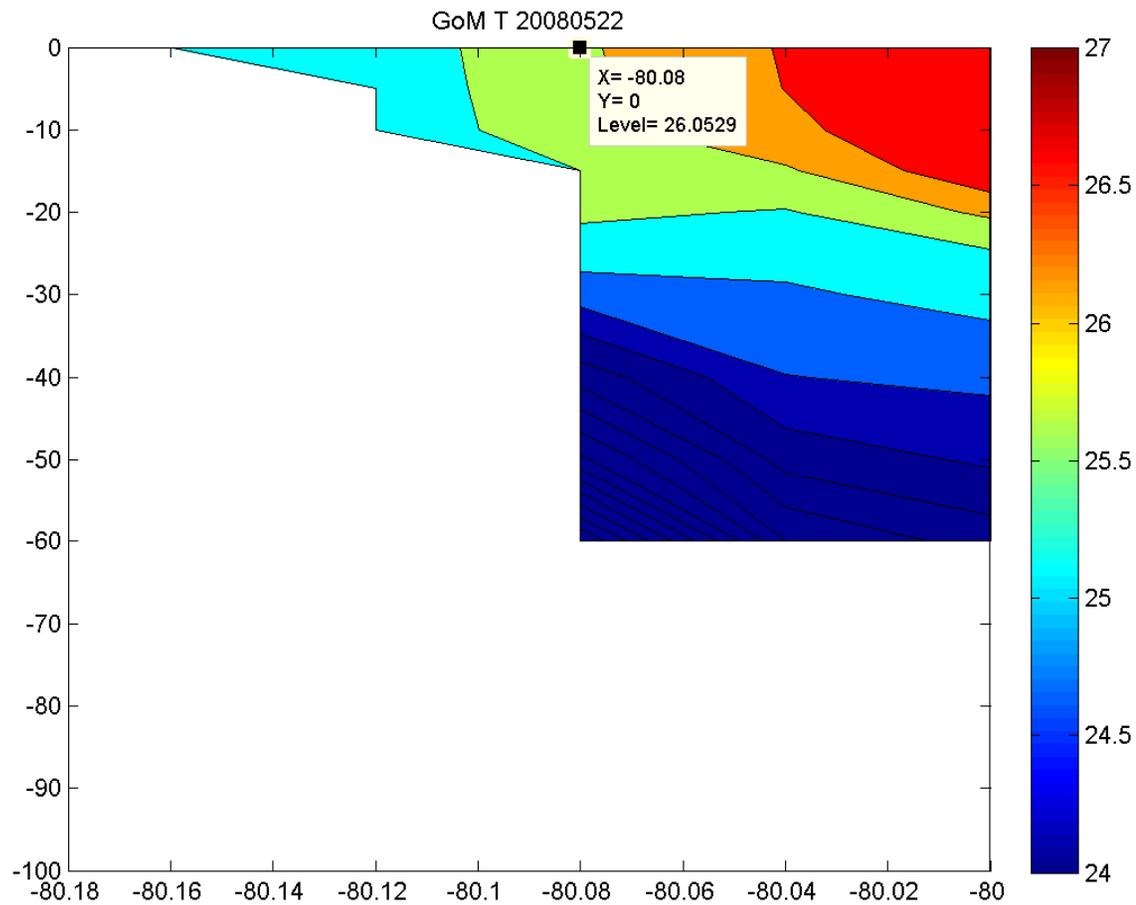


Monismith et al., 2006

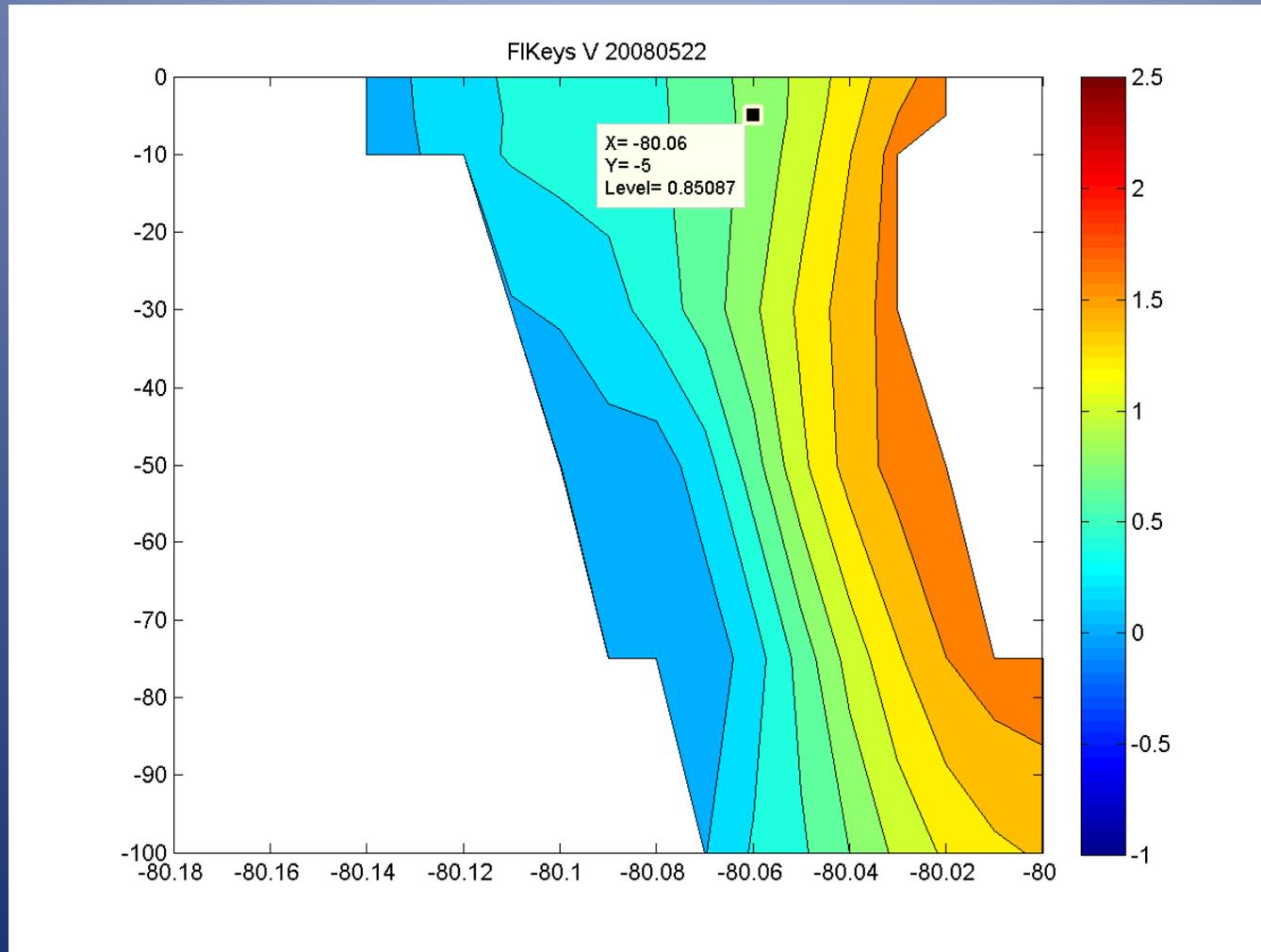
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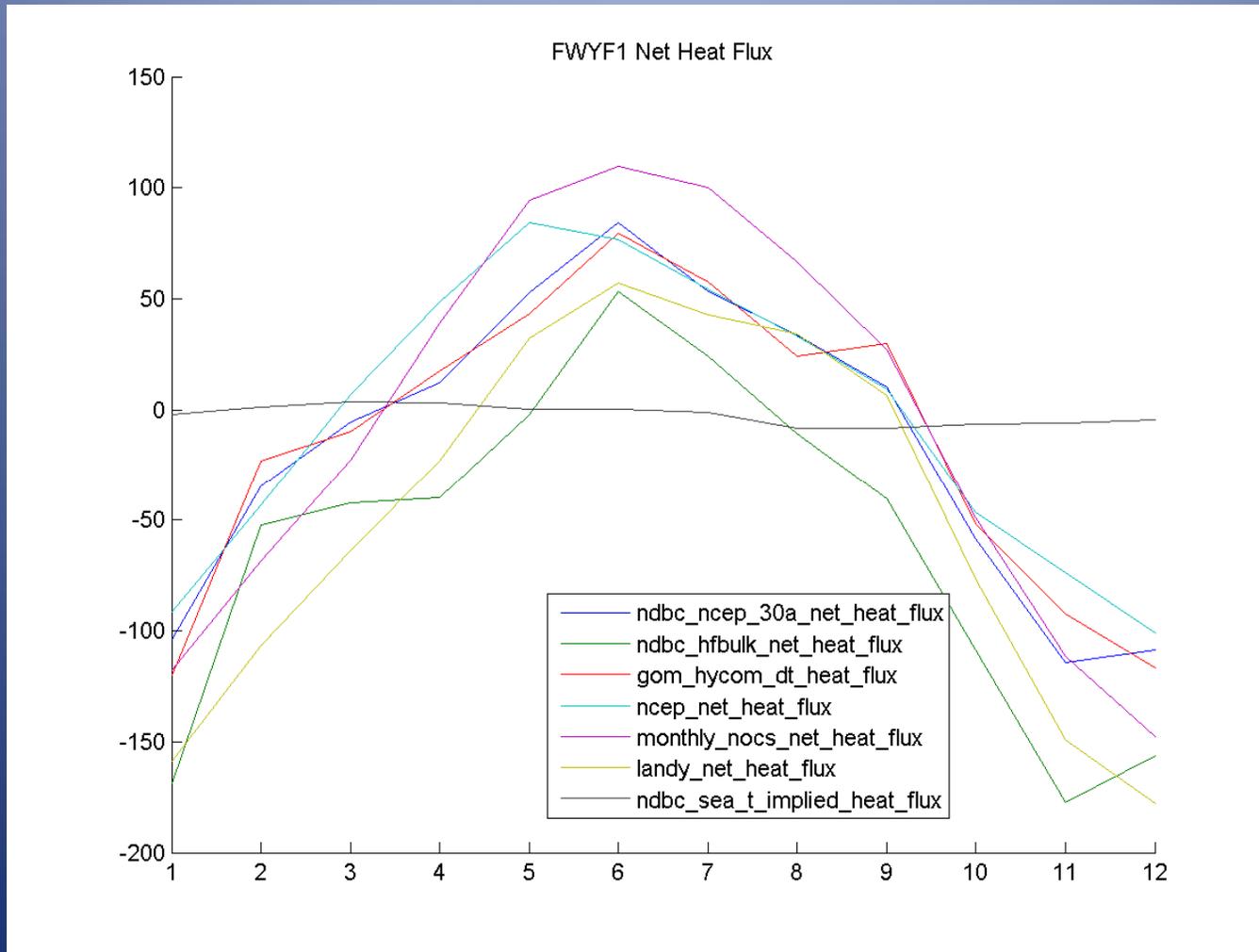
Cross-shore temperature profile – Gulf of Mexico HYCOM (May 2008)



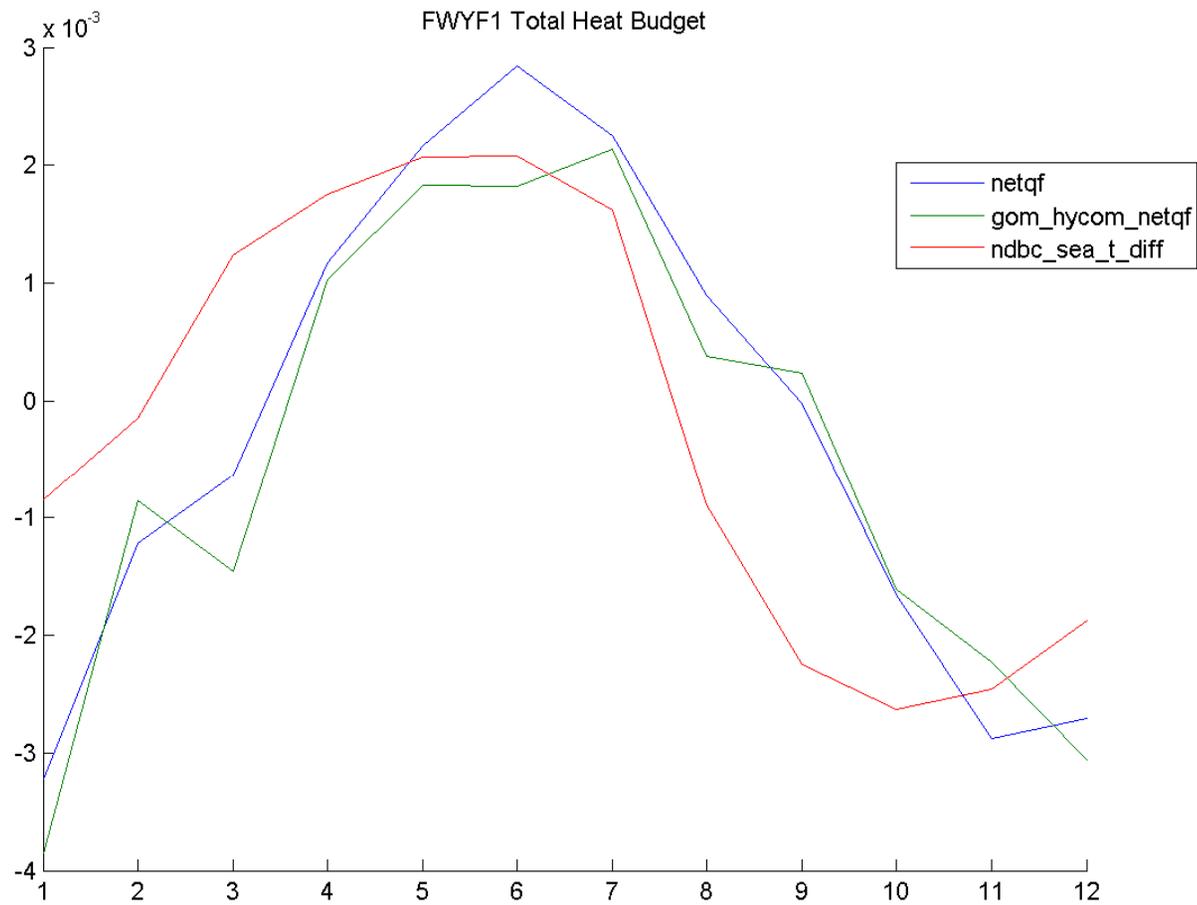
Cross-shore along-shore velocity – Florida Keys HYCOM (May 2008)



Monthly climatology from literature and current work – air-sea fluxes only



Monthly climatology – with thermal siphon



Weekly climatology – with thermal siphon

